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Mapping monthly rainfall data in Galicia (NW Spain) using inverse distances and geostatistical methods

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Abstract. In this paper, results from three different interpolation techniques based on Geostatistics (ordinary kriging, kriging with external drift and conditional simulation) and one deterministic method (inverse distances) for mapping total monthly rainfall are compared. The study data set comprised total monthly rainfall from 1998 till 2001 corresponding to a maximum of 121 meteorological stations irregularly distributed in the region of Galicia (NW Spain). Furthermore, a raster Geographic Information System (GIS) was used for spatial interpolation with a 500×500 m grid digital elevation model. Inverse distance technique was appropriate for a rapid estimation of the rainfall at the studied scale. In order to apply geostatistical interpolation techniques, a spatial dependence analysis was performed; rainfall spatial dependence was observed in 33 out of 48 months analysed, the rest of the rainfall data sets presented a random behaviour. Different values of the semivariogram parameters caused the smoothing in the maps obtained by ordinary kriging. Kriging with external drift results were according to former studies which showed the influence of topography. Conditional simulation is considered to give more realistic results; however, this consideration must be confirmed with new data.

1 Introduction

Rainfall is one of the main variables that define the climate of a region so its spatial characterisation is highly relevant. In Galicia, Northwestern Spain, the influence of topography and distance from the coast on rainfall had been described previously from annual isohyets maps obtained by traditional interpolation methods (De Uña Álvarez, 2001). Highest rainfall values observed for an average of 30 years are located on the eastern mountains and on the precoastal elevations

and the lowest ones are observed in the interior fluvial valleys. Annual rainfall amounts show a complex distribution because of the importance of the interactions between atmospheric and geographical variables (Carballeira et al., 1983; De Uña Álvarez, 2001). In this region, an increase of rainfall with elevation and a decrease of rainfall with distance to the shoreline is an expected result, because of the general circulation pattern with fronts arriving from the West, mainly from the Southwest. However, generally, a loss of regularity in the rainfall regime is observed when the distance to the shoreline increases, which may be attributed to the influence of topography at smaller scale (De Uña Álvarez, 2001).

Rainfall is a spatio-temporal intermittent phenomenon displaying large spatial and temporal variability and rain gauge networks only collect point estimates. Thus, providing an estimate of rainfall spatial distribution within an area from raingauges data frequently remains a problem of interpolation. Indeed, other methods do exist which can provide better resolution without interpolation, such as radar measurements, but nowadays only one radar is installed in our study region so that and of the area of Galicia is not covered by this device.

Traditional interpolation methods, such as inverse distances, are appropriate for a rapid estimation and they can be integrated into a Geographical Information System (GIS) as well. On the other hand, geostatistical methods require a complete previous analysis of the spatial structure of the studied phenomenon and might be non-applicable in all the occasions but they perform a better interpolation when the variable shows a spatial dependence.

Several studies have demonstrated the convenience of performing geostatistical analysis for mapping rainfall in different geographical locations (Goovaerts, 2000; Gómez Hernández et al., 2001; Militino et al., 2001), mainly at the yearly scale. Moreover, the use of a complex technique is no warranty of a better performance in a given region (Gómez Hernández et al., 2001).

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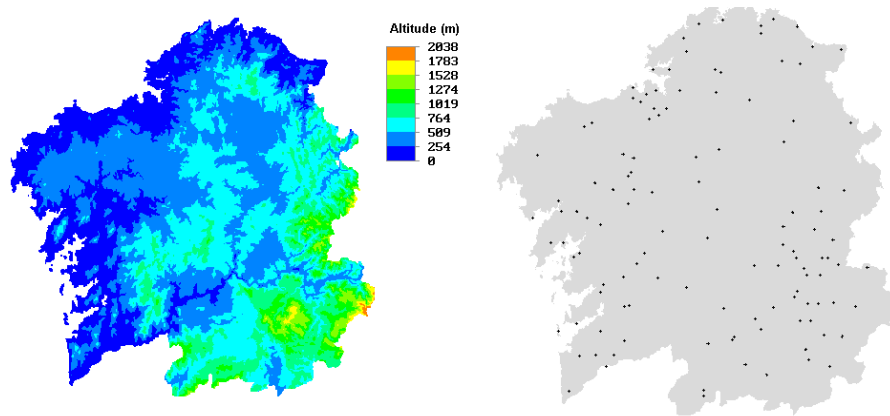


Fig. 1. Digital elevation model of Galicia and raingauge distribution.

In this paper, monthly rainfall data from Galicia (NW Spain) are interpolated using a deterministic method (inverse distances) and geostatistical methods (ordinary kriging, kriging with external drift and gaussian conditional simulation). Interpolation performances of the different algorithms are compared.

2 Data and domain

Data sets used in this work corresponded to total monthly rainfall (in millimetres) during the period from January 1998 to December 2001 and are referred to 121 rain gauges which are distributed irregularly on Galicia. In order to carry out spatial interpolation using geostatistical techniques, a complete analysis of the spatial structure of this data set was performed using GSTAT software (Pebesma, 2000) integrated in a GIS called PCRaster (Van Deursen and Wesseling, 1992).

The basis for performing interpolations by both geostatistical and inverse distance methodologies was a digital elevation model of Galicia (Thonon and Paz González, 2004) which consists of regular cells of 500 by 500 m covering an area of 29 750 km² (Fig. 1). Raingauge distribution is shown in Fig. 1 as well.

Data sets were characterised statistically; this description included the calculation of mean, median, mode, minimum, maximum, coefficient of variation, standard deviation, skewness and kurtosis. Number of data varied from 71 to 121 measurements depending on the month.

Moreover, an analysis of correlation between rainfall and altitude and distance to the shoreline was carried out in order to incorporate comprehensive altitude and distance to the shoreline data as secondary information in the interpolations.

3 Interpolation methods

3.1 Inverse distances interpolation

Inverse distances method was used as a reference for mapping monthly rainfall data: rainfall is estimated as a linear combination of several surrounding observations, with the weights being inversely proportional to the square distance between observations and the point to be estimated (Burrough and McDonnell, 1998).

3.2 Semivariogram analysis

Spatial variability was primarily evaluated through semivariogram calculation, graphing, model fitting and comparison for each variable (Burgess and Webster, 1980). To verify the spatial variability of rainfall, results were analyzed as described by Vieira (2000), based on the assumptions of stationarity in accordance with the intrinsic hypothesis. Geostatistical data treatment comprised the analysis of semivariograms and fitting of their respective model parameters. Based on the mathematical model used to fit the data, the following semivariogram parameters were defined: a) nugget effect (C_0), which is the value of the semivariogram when distance is 0; b) range of the spatial dependence (d), which is the range or distance at which semivariogram remains approximately constant, after increasing as distance increases; c) threshold ($C_0 + C_1$) which is sill value approaching the data variance.

All semivariograms were estimated using GSTAT software (Pebesma, 2000). Initially, the method used to select the parameters (nugget effect, range, sill) of the different semivariogram models was the weighted least-squares method, where the weights were the number of datapairs that provided the information needed to calculate the experimental semivariograms. Then, the cross-validation technique (Chilés and Delfiner, 1999) was used to check model performance.

Table 1. Statistical summary of the monthly rainfall of 1999 (SD = standard deviation; CV = coefficient of variation).

Date	N	Mean (mm)	Median (mm)	SD (%)	CV	Min. (mm)	Max. (mm)	Mode (mm)	Skewness	Kurtosis
January	103	121.19	114.7	58.5	0.48	34.5	357.3	81	1.66	4.21
February	105	60.53	57.5	24.9	0.41	12	158.2	49	1.04	1.94
March	109	164.24	152.0	78.7	0.48	46.1	429.6	81	0.9	0.75
April	105	159.03	155.2	56.7	0.36	20.79	352.4	195.1	0.57	0.55
May	105	123.03	117.0	42.1	0.34	46	268.3	85	1	1.15
June	101	30.98	29.9	15.2	0.49	3.8	69.3	18	0.43	−0.55
July	94	19.82	17.5	11.3	0.57	4.2	64.8	13	1.59	3.12
August	104	70.90	69.0	29.5	0.42	12.3	153	81	0.47	−0.27
September	109	232.33	231.5	84.6	0.36	84.4	472.4	124	0.4	−0.04
October	105	222.86	214.0	88.2	0.40	62.3	502	142	0.75	0.42
November	101	91.75	69.5	47.0	0.57	16.5	242.2	56	0.88	0.56
December	110	210.26	185.5	90.2	0.43	83.5	599.6	106	1.51	3.14

Because of the limited number of measured data only the omnidirectional semivariogram was computed, and hence the spatial variability is assumed to be identical in all directions. Classic criteria for calculating semivariograms were taken into account (Samper Calvete and Carrera Ramírez, 1996; Goovaerts, 1997).

3.3 Kriging and simulation

Geostatistical interpolation was performed according to the following methods:

- Ordinary kriging (OK).
- Kriging with external drift (KED).
- Gaussian conditional simulation (GS).

The complete mathematical description of these methodologies is beyond the scope of this report (Samper Calvete and Carrera Ramírez, 1996; Goovaerts, 1997; Chilés and Delfiner, 1999) so they are described only briefly.

OK is by far the most common type of kriging in practice (Webster and Oliver, 2001). Kriging interpolation methods provide each cell with a local, optimal prediction and an estimation of the error that depends on the variogram and the spatial configuration of the data (Burrough and McDonnell, 1998). Kriging is a generalized technique that allows one to account for the spatial dependence between observations, as revealed by the semivariogram, into spatial predictions (Goovaerts, 2000). The OK weights are determined such as to minimize the estimation variance (Goovaerts, 2000).

KED incorporates secondary information into the kriging system when the main and secondary variable are correlated. It is necessary to know the value of the secondary variable at all points where the primary variable is going to be estimated (Goovaerts, 1997). The secondary information is used to find

the local means of the primary variable and performs simple kriging on the corresponding residuals (Goovaerts, 2000).

Conditional simulation was performed as described by Deutsch and Journel (1992). The method used the Gaussian conditional simulation through LU (lower-upper) decomposition of the variance matrix. Conditional simulation combines the data at the observation points with the information from the variogram to compute the most likely outcomes per cell as a function of the variogram parameters. When a sufficient realization number is used, mean values obtained by simulation are expected to be similar to those obtained by kriging. Conditional simulation was performed as described by Deutsch and Journel (1992). When a sufficient realization number is used, mean values obtained by simulation are expected to be similar to those obtained by kriging. This method does not provide estimation error maps. For conditional simulation analysis, 100 different realizations were run. From these data, a mean value was computed as well as standard deviation.

4 Results and discussion

4.1 Statistical analysis

Mean annual rainfall during the study period varied from 1214.3 mm in 1998 to 1772.5 mm in 2000; the wettest month was March 2001 (467.4 mm) and the driest one was August 1998 (6.6 mm). Table 1 shows the results for the statistical analysis of monthly rainfall during 1999, as an example.

Coefficients of variation (defined as the ratio of the standard deviation to the mean) of monthly rainfall values, which varied between 20% and 146%, showed the spatial heterogeneity of the rainfall in Galicia; this variability was higher in the dry season, from June to August (see Table 1 and Mirás Avalos, 2003).

Table 2. Rainfall mean values estimated by the different methodologies used in this study. Data appear in millimetres.

Month	Measured	Inverse distance	OK	KED altitude	KED distance to the shoreline	GS
January 1998	143.2	142.8	144.2	122.4	138.9	158.9
June 1998	31.4	31.1	30.8	32.2	31.0	33.3
July 1998	43.8	44.9	45.8	43.5	45.7	50.6
August 1998	6.6	6.9	7.3	7.8	7.3	8.9
November 1998	82.1	82.0	81.5	87.1	82.2	86.6
December 1998	121.4	120.8	120.2	112.8	118.2	132.6
January 1999	121.2	121.2	120.7	116.1	117.8	133.1
February 1999	60.5	60.2	60.1	60.7	59.1	66.1
March 1999	164.2	166.3	163.3	172.4	166.0	183.6
April 1999	159.0	160.3	158.6	146.0	156.6	177.4
May 1999	123.0	123.3	124.1	121.1	123.3	136.4
July 1999	19.8	20.2	20.0	19.8	20.1	22.6
August 1999	70.9	70.5	70.6	66.8	70.0	78.4
September 1999	232.3	233.3	230.6	221.5	229.1	260.8
October 1999	222.9	222.8	225.2	203.9	223.6	248.1
November 1999	81.8	82.3	84.0	86.8	83.5	90.1
December 1999	210.3	206.9	202.3	213.8	204.2	226.7
March 2000	29.7	30.4	30.8	31.0	30.8	34.5
April 2000	314.8	312.3	306.9	313.9	306.2	342.3
May 2000	96.0	95.9	95.2	96.7	95.3	106.5
June 2000	16.3	16.6	17.0	18.1	16.8	18.6
July 2000	60.9	60.2	62.3	62.4	61.6	68.7
August 2000	47.3	46.0	43.7	44.7	43.7	48.2
September 2000	79.7	79.2	79.1	73.1	77.0	87.0
October 2000	173.2	172.2	171.5	161.4	168.5	189.1
November 2000	410.3	404.7	398.4	397.3	394.7	439.9
December 2000	445.4	451.2	458.9	461.1	451.9	508.1
January 2001	368.1	370.2	365.0	373.9	365.4	404.4
March 2001	467.4	465.1	459.7	476.8	460.4	511.6
April 2001	82.5	81.5	81.7	84.7	80.9	88.3
July 2001	82.1	82.8	82.8	86.4	83.4	92.0
August 2001	56.6	57.0	56.0	59.3	56.2	60.8
October 2001	226.9	224.1	216.5	224.7	216.4	251.2

Skewness and kurtosis coefficient values were close to those of a standard Gaussian variable. From this test, it was assumed that monthly rainfall data followed a Gaussian distribution, which is a requisite for standard semivariogram analysis.

On the other hand, by linear regression analysis it was found that correlation between altitude and rainfall ranged from 0.04 to 0.55, whereas correlation between rainfall and distance to the shoreline varied from 0.004 to 0.71. The use of this comprehensive secondary information into the mapping of rainfall was only worth accounting when correlation coefficients were statistically significant.

4.2 Spatial dependence analysis

Different theoretical functions (spherical, exponential, linear and gaussian) with variable nugget effects depending on the analyzed month were fitted to experimental data. Rainfall

spatial dependence was observed in 33 out of 48 months analyzed; the rest of the rainfall data sets presented no spatial structure and were modelled by a pure nugget effect. The most frequently used model was the exponential which was fitted to 18 monthly data series; spherical model was fitted to 13 data series, gaussian and linear models were fitted to one data set each. An example of experimental semivariogram and its fitted model is shown in Fig. 2, each number in the figure represents the number of pairs for each lag. Range and nugget effect of the fitted models were variable fluctuating, the former, between 2.99 and 59.87 km and, the latter, between 0 and 64.02% of the sill value. Generally, spatial dependence ranges were higher for the spherical models than for the exponential ones; ranges were lower than 10 km in 8 occasions, were between 10 and 30 km in 14 months and were higher than 30 km 11 times. Due to the short spatial dependence range, kriging interpolation results showed high

kriging estimation errors as the distance to the rain gauges increased. Nugget effects were nule or lower than 10% in 15 months; between 10 and 25% in seven months and higher than 25% in 11 months. Therefore, half of the studied data series showed a high spatial dependence. Cross-validation showed that the theoretical structures described the spatial dependence of rainfall adequately (data not shown; Mirás Avalos, 2003).

4.3 Inverse distances

Inverse distance technique was the only method that could be used for mapping monthly rainfall in Galicia in all the studied months. Maps obtained by this technique showed a division in more or less distinct areas and, within them, usually punctual areas corresponding to high or low rainfall values were observed, which resulted in a discontinuous appearance of the maps (Fig. 3). Generally, differences between mean values measured at the gauges and those obtained by inverse distances method were lower than ± 10 mm (Table 2).

4.4 Ordinary kriging (OK)

Maps of mean monthly rainfall obtained by this method showed distribution patterns similar to those obtained by inverse distances (Fig. 3). However, these maps were too smooth for reproducing measured maximum and minimum data; moreover, error maps tended to show a high and uniform uncertainty pattern (Fig. 4). Thus, OK offered both smoothed maps and high kriging errors. In fact, estimated extreme values showed a lower oscillation ratio than those which were measured, but sometimes they were not very different. Overall, OK tended to underestimate rainfall mean values (Table 2). Kriging errors were high in most of the study area except in those areas located nearby the rain gauges. In fact, due to the small range of spatial dependence, generally in the order of 10 km, kriging errors with small values were limited to a close neighbourhood around the rain gauges.

4.5 Kriging with external drift (KED)

Additional information (altitude and distance to the shoreline) was used in order to enhance the previous estimation by using the KED technique. A noticeable influence of topography over the rainfall estimated values was observed (Fig. 3). The influence of altitude over rainfall estimated values is more obvious when spatial dependence range is lower than 10 km; however, when distance to the shoreline was used as secondary information, this technique gave maps similar to those obtained by ordinary kriging.

KED tended to underestimate monthly rainfall mean values up to 15% (Table 2). Results using altitude as secondary information showed differences regarding those of ordinary kriging and inverse distances. Moreover, KED estimation uncertainty showed a similar pattern to that obtained by OK

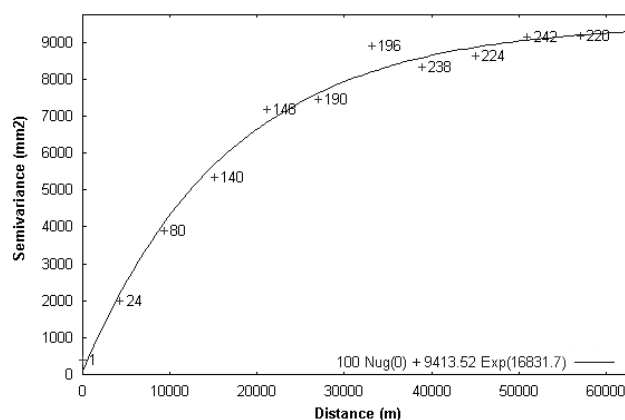


Fig. 2. Example of experimental variogram and its fitted model, April 2000.

and they showed similar values, as well (data not shown; Mirás Avalos, 2003).

4.6 Gaussian conditional simulation (GS)

This technique gave average maps that presented a variability which made them to seem real, reproducing, simultaneously, measured values (Fig. 3). Generally, coefficients of variation of each 100 simulation set were lower than 25%. This methodology seemed to overestimate the rainfall mean values between 2% and 35% (Table 2).

Conditional simulation interpolation reproduced monthly rainfall minimum and maximum values when nugget effect was nule or low. However when nugget effects were high, oscillation ranges of simulated values were lower than those registered at the rain gauge network.

The fact that montly mean rainfall values obtained by conditional simulation were systematically higher than those registered, might be a good approximation to reality, because of the irregular distribution in altitude of the rain gauges used for this study. In other words, higher values of rainfall are expected in mountain areas where the rainfall network is scarce, thus conditional simulation results could be a more accurate approximation to real rainfall over Galicia than mean values obtained from measurement network.

4.7 Comparison of kriging and simulation results

Like the inverse square distance method, geostatistical interpolation aims at estimating the unknown rainfall depth at the unsampled location as a linear combination of neighbouring observations. The main advantage of kriging is to provide kriging estimation error maps. OK main drawbacks in this study were the smoothing of the estimation maps and the high values of the kriging errors.

KED yields maps whose trends reflect the variability of the secondary attribute but it is not proof that those trends

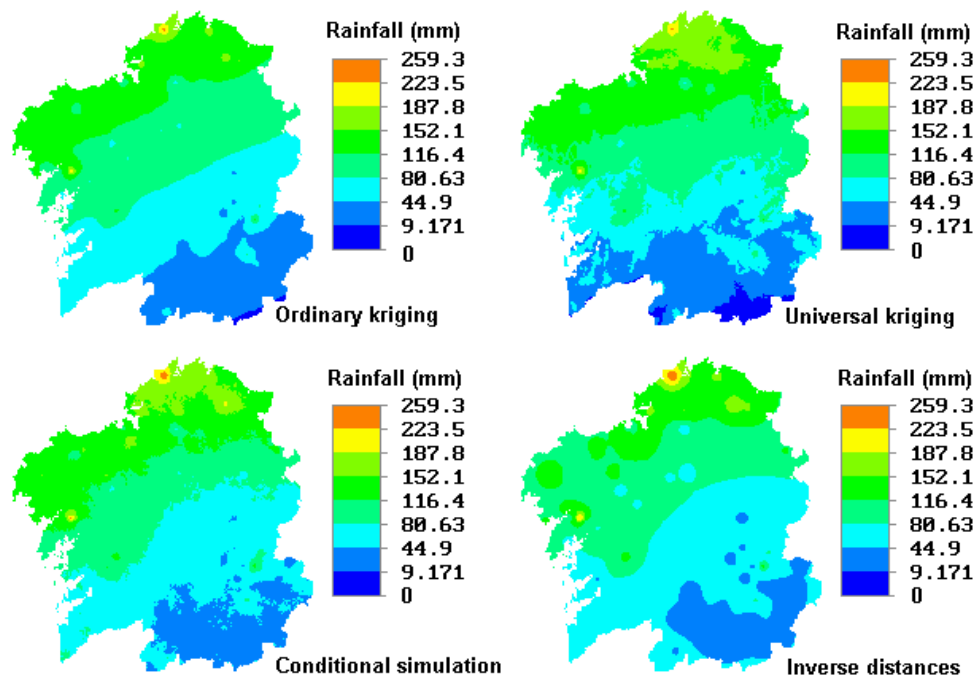


Fig. 3. Rainfall distribution for November 1999 interpolated by different techniques.

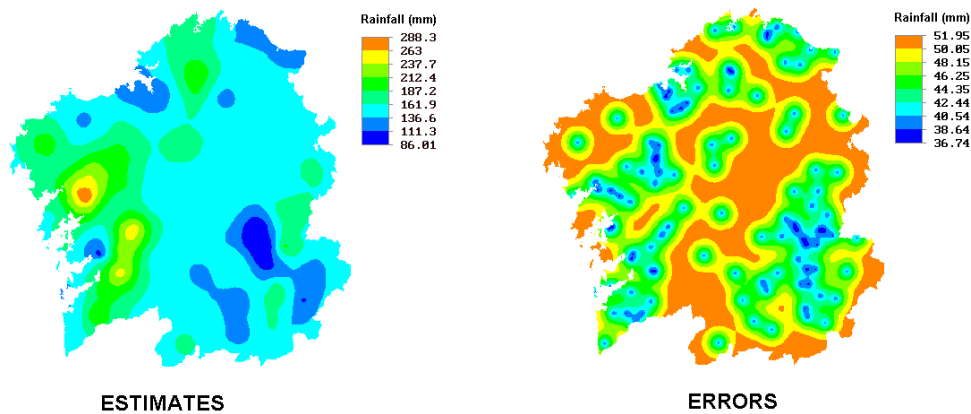


Fig. 4. Rainfall distribution for April 1999 interpolated by ordinary kriging and estimation errors for the same data set.

necessarily pertain to the study variable (Deutsch and Journel, 1992). The relation between elevation and rainfall is assessed locally, which allows one to account for changes in correlation across the study area (Goovaerts, 2000). In this exercise, KED maps showed a high influence of the altitude when semivariogram ranges were short. This effect of the altitude on the rainfall was previously described by other authors. For this reason and taking into account the topographic characteristics of Galicia, interpolation performed using altitude as external drift was more plausible than the one performed using distance to the shoreline.

Kriging and conditional simulation are geostatistical techniques that have the advantage of optimizing an interpola-

tion variable. Nevertheless, the aims of these techniques are different: kriging tries to minimize the estimate variance error while conditional simulation tries to reproduce the spatial variability of the studied phenomenon.

Conditional simulation has been considered to be a method for optimizing kriging results. Some authors (Samper Calvete and Carrera Ramírez, 1996; Chilés and Delfiner, 1999) suggested to combine conditional simulation results with those of kriging in order to achieve the advantage of knowing the estimation uncertainty. In this study, GS provided a better description of the maximum and minimum rainfall values than that offered by other methods. Monthly rainfall mean values were higher than those recorded, what could be

approximate to the real values due to the irregular distribution in altitude of the rain gauges used.

5 Conclusions

Structural analysis of the study data set showed that not a single model could be fitted to all of the experimental data sets, however, different models were fitted to 33 out of 48 data sets. Therefore, the fitted models depended on the pluviometric characteristics occurred in each month.

Inverse distance interpolation was the only method that could be used for mapping Galicia monthly rainfall in all the studied months. It offered discontinuous maps and proved to be appropriate for a rapid estimation of rainfall.

OK gave maps that showed similar patterns to those obtained by inverse distances but it provided estimation error maps as well. OK estimated maps were too smooth for reproducing original maximum and minimum data; moreover, kriging error maps tended to show a high and uniform uncertainty pattern.

KED offered reasonable results; however, drift approach seemed to be arbitrary. Rainfall spatial distributions and monthly rainfall mean values obtained by the two different KED approaches, usually, differed notably. Generally, when distance to the coast was used as secondary information, isohyets showed a notorious smoothing, thus, the maps were similar to those obtained by ordinary kriging. Taking into account the topographic characteristics, interpolation performed using altitude as external drift was more plausible.

Finally, GS gave average maps that seemed to fit very well to what occurred in reality. This technique reproduced monthly rainfall maximum and minimum values when nugget effect was null or low; in contrast, when this nugget effect was high, maximum and minimum value oscillation ratio was lower than that registered at the gauges.

Taking into account that mountainous areas, with higher precipitations, are not well represented in the sample population, conditional simulation is supposed to offer more realistic results; nevertheless, this consideration must be confirmed with new measured data.

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